

FLOW EXTENSION FOR SELECTED GAGES IN
THE WESTERN U.S.

Phase II of II – Streamflow Record Extension in the
Grand Mesa-Uncompahgre –Gunnison National Forest,
Colorado

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PHASE II of II

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Streamflow Record Extension in the Grand Mesa-Uncompahgre-Gunnison National Forest,
Colorado

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Introduction

This text reports on the extension of stream gage records in the Grand Mesa-Uncompahgre-Gunnison (GMUG) National Forest in Colorado, such that each extended record covered at least water years 1970-99. Extensions were performed by relating each existing record to another gage in the Forest covering this period. This report discusses the selection of these pairs of gages, data preparation, and the extension method.

All flows considered herein are daily flows as reported by the USGS NWIS-W, and all extended records are presented as sets of daily flows. The word *extension* is used to indicate not only an increase in the overall time span of a gage, but also the estimation of missing dates within that period. The gage record to be extended is called the *short* record, while the one used to extend it is known as the *base* record.

Records extended

The gages in the GMUG National Forest for which daily flow records were extended are listed in the Appendix. Of the 77 gages in the Forest, four had continuous records for the 1970-1999 period and so were used as base stations to extend the other records. Of the other 73, 14 were found to be unsuitable for extension using the available base records. As a result, the record at each of 59 gages was extended for a period corresponding to the extent of its base record, which in all cases covered at least the 1970-1999 water years.

The suitability for extension of each short record was assessed on the basis of its relationship to that base record having the greatest correlation to the short record's daily flow values during the *concurrent period*, the set of days on which both gages recorded discharges. The assessment employed both the correlation coefficient r and scatterplots of the concurrent daily flows at the two gages. Screening rules were based primarily on experience. They may be summarized as:

1. Reject records with a concurrent period shorter than four years, to insure suitable representation of the statistical relationship.
2. Reject all records with a concurrent-period correlation below 0.75.
3. Reject records with unusual structure in a scatterplot of the short vs. base record. This is generally a result of flow regulation, and such cases might be rehabilitated by removal of a portion of the record.
4. Accept all remaining records with a concurrent-period correlation above 0.85.
5. The suitability of the remaining records is assessed on the basis of
 - a. the overall concurrent-period correlation,
 - b. length of the concurrent period,
 - c. homoscedasticity (uniformity of variance) or unusual structure in the scatterplot,
 - d. the number of flows recorded as zero, and

- e. the number of flows so low that the gage was imprecise and possibly inaccurate.

Five records (9, 32, 33, 55, 64) were rejected by rule 1 and sixteen others (5, 6, 7, 27, 41, 42, 48, 60, 61, 62, 65, 67, 70, 72, 73, 75) by rule 2. None of the remaining records were rejected by rule 3. Rule 4 passed 42 of the remaining 52 records. Of the remaining 10 records subject to judgement (rule 5), five (8, 13, 16, 39, 40) were accepted and five (4, 51, 54, 56, 74) rejected, for a total of 47 extended short records.

The causes of low correlation and the criteria of rule 5 are examined in greater detail later in this text.

Data Preparation

All flows were downloaded from the USGS NWIS-W data set at <http://waterdata.usgs.gov/nwis-w/US/>. Estimated flows were accepted as given

The only data preparation was the handling of zero flow data. In the extension process, all flow data were subject to transformations (such as the logarithm) which cannot be used with zero flows. Three of the records which were extended had zero data: 2, 40, and 66, with zeroes comprising 0.2%, 0.4%, and 0.04% of the data, respectively.

Handling of these flows followed a procedure described by Stedinger et al. (1993). Zero flows may be regarded as those with magnitudes below the detection threshold of the gage, and thus below all the measured values. The set of zero dates was assigned a set of exceedance probabilities according to the Weibull plotting position, $i/(n+1)$, where i is the rank of the flow and n is the number of dates. The flow magnitudes for this set were taken as the lower tail of a lognormal distribution, such that the highest flow in the set equaled the lowest nonzero recorded flow, and the mean of the overall lognormal distribution equaled the mean log flow of the set formed by the union of the observations with the set of zero dates.

While the values assumed for the zero-flow dates generally matter little in application - the difference between zero and the lower detection limit being insignificant - the values used in the extension can make a difference owing to transformations. The objective of filling in these data therefore was to make them as consistent as possible with the known data while falling below the known data.

Extension technique

Each extended record is the union of two sets of flows: (A) the known flows for that gage and (B) flows generated for dates covered by the base stream gage, but not appearing in the short record. The calculation of short-record flows from base station flows proceeds from the generation of an isomorphic, continuous, monotonic curve relating the records.

MOVE.1 with scaled power transformations

The relationships used to extend the records from the base stations were generated using an extension of the technique described by Moog et al. (1999). In this technique, the data for each record are subjected to a scaled power transformation:

$$Q' \equiv \begin{cases} \frac{Q^\lambda - 1}{\lambda} & \text{if } \lambda \neq 0 \\ \ln Q & \text{if } \lambda = 0 \end{cases} \quad (1)$$

where Q is the flow rate and λ is an exponent chosen by a maximum likelihood criterion. Note that the logarithmic transformation is simply a special case of the more general scaled power transformation, which can linearize relationships which would be curved in log-log space. The goal of the transformation is to condition the data so that the relationship between the records is more linear and homoscedastic when a linear fit is calculated.

The MOVE.1 equation (Hirsch, 1982) is used to relate the transformed data linearly such that the estimated mean and variance for the concurrent data equal the mean and variance of the measured data. The equation takes the form

$$\hat{y}_E = m(y_C) + \frac{S(y_C)}{S(x_C)}(x_E - m(x_C)) \quad (2)$$

where $\hat{y}$ indicates a set of estimated flows, y a set of short-record flow measurements, x a set of base record flows, subscript C the concurrent period, and subscript E the rest of the extension period. The functions m and S are the sample mean and standard deviation. In the current application, x and y were transformations of the flow rate using Eq. (1). This equation is the same as that for linear least-squares regression with the coefficient r of the last term replaced by unity. Thus the slope of the line is greater than that produced by least-squares linear regression, equal to that of a record perfectly correlated to the base record. The result is greater variance in the extended record than would be produced by least-squares linear regression.

After application of Eq. (2), the results are retransformed - that is, back to flows in cfs - by inverting Eq. (1). The retransformation may produce complex numbers for some values of λ ; the values are thus chosen conservatively following the guidelines of Moog et al. (1999).

Modifications to MOVE.1 for retransformation bias

For generation of streamflow records, the Moog et al. (1999) technique has one shortcoming: retransformation bias. While the mean and variance of the measured and estimated data are equal for the transformed variables, they differ after retransformation to cfs. This was not a problem in the 1999 paper, for which only the shape of the flow duration curve mattered, not the actual magnitude of flow. While there may be an argument that matching in the transformed space is preferable, hydrologists would generally prefer a formula which produces the same mean flow in cfs as that of the measured data (during the concurrent period).

For this reason, the MOVE.1 technique was modified to remove the retransformation bias. While there are a number of bias correction techniques for methods, like linear least-squares regression, which estimate the line of means, we knew of none available for MOVE.1. However, the desired modification is easily defined; the new MOVE should match the measured mean and variance in the original space (e.g., cfs) rather than the transformed space, while remaining linear in transformed space.

The two constraints (matching mean and variance) lead to two equations in two unknowns (slope and intercept of the line), as follows. The desired linear equation in transformed space is

$$\hat{y}'_c = a + bx'_c \quad (3)$$

where the primes indicate transformed flows, and a and b are the unknown parameters which we wish to find. To match the mean and variance in the retransformed flows (cfs), we require:

$$\sum \hat{y}_c = \sum y_c \quad (4)$$

and

$$\sum (\hat{y}_c - m(\hat{y}_c))^2 = \sum (y_c - m(y_c))^2 \quad (5)$$

where summations are taken over the concurrent data. Using Eq. (4), Eq. (5) simplifies to

$$\sum \hat{y}_c^2 = \sum y_c^2 \quad (6)$$

Eq. (3) may be written in terms of retransformed flows by using Eq. (1):

$$\hat{y}_c = \begin{cases} (\lambda_s(a + bx'_c) + 1)^{1/\lambda_s} & \text{if } \lambda_s \neq 0 \\ \exp(a + bx'_c) & \text{if } \lambda_s = 0 \end{cases} \quad (7)$$

where λ_s is the exponent in Eq. (1) for the short record. Eq. (7) may be substituted into Eqs. (4) and (6), leading to two equations in two unknowns, a and b , implying a unique solution for the line which meets the constraints.

Solution of the modified, hybrid MOVE.1 equations (MOVE.1a)

A closed-form solution of the equations was not apparent, and there are no general techniques for numerical root-finding in two dimensions. However, it was discovered that the equations could be solved without difficulty in nearly all cases (only one of over 100 cases proving troublesome), owing to two favorable conditions. The MOVE.1 solution provides good initial guesses for the slope and intercept, and the equations are largely decoupled, in that the variance is affected primarily by the slope, and the mean by the intercept.

Therefore, the successful numerical scheme involved alternately solving the mean-matching equation (4) for the intercept, and the variance-matching equation (6) for the slope. Using the method of successive approximations, each solution was adopted as the slope or intercept for the next calculation, to convergence. Solution of the root for each equation was performed using a polynomial interpolation function (*uniroot*) in the statistical computing language S-plus. (This function requires pre-bracketing the root, which was performed by stepping away from the initial guess, and reversing with smaller steps if one step led to imaginary retransformed flows. This adds complications because for some values of λ_s and x'_c , the solution to Eq. (7) is complex so that these values must be avoided.)

Examination of the results using matching in the original space showed great improvement in estimation of high flows, but degraded estimation of low flows. This occurred because low flows have much less relative impact on the variance and mean in the original space than in transformed (e.g., logarithmic) space.

This fact, however, led to its own solution, in that it is possible to improve low-flow estimates while retaining accurate overall statistical matching. A hybrid scheme was adopted, as illustrated by hypothetical Figures 1 and 2. Figure 1 shows an example MOVE.1 curve (for which mean and variance are matched in transformed space) and a possible corresponding curve labeled MOVE.1b, based on matching in untransformed space. (The MOVE.1b slope may be lesser than that of MOVE.1, and actual curves are usually much closer.) As depicted in Figure 2, the hybrid scheme uses the MOVE.1b curve for flows above the intersection point Q_c of the curves, and MOVE.1 below. The estimated mean and variance of the concurrent flows then differ slightly from observations. Judging that matching the mean is of primary importance, the slope of the high-flow branch is then tweaked to produce exact matching of the mean, as shown to an exaggerated degree by the arrow in Figure 2. In practice, tweaking had little effect on variance, which remained much closer to observations than under MOVE.1. The resulting modified MOVE.1 method is labeled MOVE.1a in Figure 2.

Discussion

Each figure referred to in this discussion contains a plot showing flows in the short record for the period concurrent with the base record. Included are both the measured record and that which would have been estimated for the concurrent period by the extension curve.

Three types of plot are presented:

- flow duration curve, showing flow magnitude (log scale) vs. exceedance probability;
- hydrograph, showing flow versus date, covering the first six years of the concurrent period (chosen arbitrarily, as more years make the plots hard to read), with both log and linear axes used for flows; and
- scatterplot, plotting each short record flow (y axis) against the base record flow (x axis) on the same date. Note that the estimated flows form the extension curve.

The term "MOVE.1a" in the legend indicates the data were generated using the modified, hybrid MOVE method, while "MOVE.1" refers to the original method.

Typical results

Figure 3 shows a flow duration curve for a pair of gages in the San Juan National Forest. This pair has a fairly typical correlation of 0.95. The scatterplot for this pair is in Figure 4, and a hydrograph is in Figure 5. The important aspect of the hydrograph is that the frequencies of low and high flows are about right. Because flows at the short-record gage are correlated to those at the base gage, the temporal behavior of the estimated record tracks observations to a degree, but such tracking is not an explicit objective of the extension technique.

An illustration of the MOVE modifications is provided by flow duration curves for another pair of gages in Figure 6, showing both the measured flows and those estimated by the modified MOVE method, MOVE.1a. Because this station pair is highly correlated, with $r = 0.97$, one would expect a close match between the curves. Nonetheless, the MOVE.1 estimates

fall above the observations, as seen in Figure 7, which shows the highest 10% of measured discharges, plus those estimated using MOVE.1 and the modified method, MOVE.1a. The improvement with MOVE.1a is clear. The improvement stems from the fact that the statistical matching is performed in terms of untransformed flow rates, which are more sensitive to high flows.

The following table compares the concurrent-period statistics of the two short records discussed above:

Record number	Mean, cfs			Standard Deviation, cfs		
	Measured	Estimated MOVE.1	Estimated MOVE.1a	Measured	Estimated MOVE.1	Estimated MOVE.1a
11	31.1	34.1	31.1	54.9	68.9	58.6
25	107.1	111.1	107.1	177.2	194.5	182.2

Flow extension objectives and limitations

The modification of the extension technique raised issues concerning the objectives of the extension. The main distinction to be made is between estimation of flows on specific days, and generation of a representative data set - i.e., a flow duration curve. These are fundamentally different objectives.

When estimating a flow on a specific day, we recognize that for any specific flow rate in the base stream, the short record exhibits a range of possible flows - hence the scatter in the scatterplot. This scatter lies about a "line of means," which we commonly estimate using linear least-squares regression; points on this line represent our estimates of the mean short-record flow for each flow in the base stream. An estimate of the flow on a given day will be the point on the line corresponding to the measured base record flow, perhaps with a confidence interval reflecting the scatter about this line. In using this point estimate, we are in essence discarding the scatter as random noise about the true relationship between the streams.

However, the true relationship in fact does include this scatter. That is, this variation about the line of means is real and should be reflected in any representative data set we synthesize for the short record. Discarding this variation, as in least-squares regression, leads to a data set with insufficient variance (Hirsch, 1982). The objective of the MOVE family of techniques is to find a line which produces sufficient variance in the synthesized record; in MOVE.1 "sufficient variance" is that which produces the same variance during the concurrent period as that which is measured.

If the objective is to produce a flow duration curve, another important consideration is the type of flows of interest: high, low, mean, those producing cumulative transport, etc. There is some tradeoff in adopting one technique over another, though the modified MOVE.1 method described above greatly reduces this tradeoff, and should make it small for well-matched pairs of stations.

The caveat that the extended data sets do not predict flows on specific days points to other limitations. Strictly, flows in the extended data sets should not be associated with specific

dates, meaning they cannot be used to estimate serial correlation or statistics for time periods shorter than the full extension period. If monthly statistics are desired, the technique could be applied separately to each month. Actually, serial correlation is captured to a degree by following the base station flows; presumably, better, directed techniques could be employed, but the extended data sets might serve as suitable estimates of serial correlation or monthly statistics.

References

- Hirsch, R.M., A comparison of four streamflow record extension techniques, *Water Resour. Res.*, 18, 1081-1088, 1982.
- Moog, D.B., P.J. Whiting, and R.B. Thomas, 1999, "Streamflow record extension using power transformations and application to sediment transport," *Water Resour. Res.*, 35(1), 243-254.
- Stedinger, J.R., R.M. Vogel, and E. Foufoula-Georgiou, Frequency analysis of extreme events, ch. 18 in *Handbook of Hydrology*, D.R. Maidment, ed., McGraw-Hill, New York, 1993.

Appendix. Record and gage numbers, base gages, and record periods for the GMUG National Forest.

Key:

Bold indicates base gage.

Italics indicate station not recommended for base gage extension.

= base gage or best base gage candidate.

Record Number	Gage Number	Gage Name	Base Record Number	Correl. To Base Record	Concurr. per. Extended per.
1	8216500	Willow Creek At Creede, Co.	34	0.94	1951-82 1938-99
2	8224500	Kerber C At Ashley Ranch, Nr Villa Grove, Co.	71	0.85	1942-82,99 1911-12,23-26, 31-34, 36-99
3	9092500	Beaver Creek Near Rifle, Co.	34	0.92	1953-82 1938-99
4	<i>9093000</i>	<i>Parachute Creek Near Parachute, Co.</i>	<i>71</i>	<i>0.75</i>	<i>1949-54,65-70,75-86 1911-12,31-34,42-99</i>
5	<i>9095800</i>	<i>Plateau Creek Near Heiberger, Co.</i>	<i>71</i>	<i>0.71</i>	<i>1958-64 1911-12,31-34,42-99</i>
6	<i>9096800</i>	<i>Buzzard Creek Bl Owens Creek, Nr Heiberger, Co.</i>	<i>71</i>	<i>0.7</i>	<i>1956-70 1911-12,31-34,42-99</i>
7	<i>9097500</i>	<i>Buzzard Creek Near Collbran, Co.</i>	<i>71</i>	<i>0.63</i>	<i>1931-34,42-81 1911-12,22-99</i>
8	9097600	Brush Creek Near Collbran, Co.	71	0.84	1956-67 1911-12,31-34,42-99
9	<i>9098500</i>	<i>Atkinson Creek Near Collbran, Co.</i>	<i>52</i>	<i>0.71</i>	<i>1953-55 1939-99</i>
10	9099000	East Fork Big Creek Near Collbran, Co.	71	0.92	1951-55 1911-12,31-34,42-99

11	9099500	Big Creek At Upper Station, Near Collbran, Co.	52	0.92	1946-56 1939-99
12	9100500	Cottonwood Creek At Upper Sta, Near Molina, Co.	52	0.91	1945-57 1939-99
13	9104500	Mesa Creek Near Mesa, Co.	52	0.83	1941-60 1939-99
14	9107000	Taylor River At Taylor Park, Co.	57	0.95	1929-34 1959-99
15	9107500	Texas Creek At Taylor Park	34	0.93	1930-34,38-99
16	9108000	Willow Creek At Taylor Park, Co.	71	0.79	1931-34 1911-12, 29-34, 42-99
17	9110500	East River Near Crested Butte	34	0.96	1940-51 1938-99
18	9111000	Coal Creek Near Crested Butte, Co.	71	0.86	1942-46 1911-12, 31-34, 42-99
19	9111500	Slate River Near Crested Butte, Co.	71	0.92	1942-51, 94-99 1911-12, 31-34, 40-99
20	9112000	Cement Creek Near Crested B	34	0.94	1940-51 1911-13, 38-99
21	9112200	East River Bl Cement Creek	34	0.94	1964-72,80- 81,94-99 1938-99
22	9113000	Castle Creek Near Baldwin	34	0.97	1945-50 1938-99
23	9113100	Castle Creek Above Mouth Near Baldwin, Co	57	0.95	1993-98 1959-99
24	9113300	Ohio Creek At Baldwin, Co.	71	0.92	1959-70 1911-12, 31-34, 42-99
25	9115500	Tomichi Creek At Sargents, Co.	71	0.86	1942-72, 93-99 1911-12, 17-22, 31-34, 38-99
26	9118000	Quartz Creek Near Ohio City, Co.	34	0.92	1938-50,60-70 1938-99

27	9118450	<i>Cochetopa Creek Below Rock Creek Nr Parlin, Co.</i>	71	0.74	1981-99 1911-12, 31-34, 42-99
28	9121500	Cebolla Creek Near Lake City, Co.	34	0.87	1946-54 1938-99
29	9122000	Cebolla Creek At Powderhorn, Co.	71	0.86	1942-55 1911-12, 31-34, 38-99
30	9122500	Soap Creek Near Sapinero, Co.	71	0.89	1955-66 1911-12, 31-34, 42-99
31	9123400	Lake Fork Below Mill Gulch Near Lake City, Co.	34	0.97	1982-86 1938-99
32	9123500	<i>Lake Fork At Lake City, Co.</i>	71	0.92	1932-34 1911-12, 18-24, 31-37, 42-99
33	9124000	<i>Henson Creek At Lake City, Co.</i>	71	0.94	1932-34 1911-12, 18-19, 31-37, 42-99
34	9124500	Lake Fork At Gateview, Co.	NA	NA	1938-99
35	9125000	Curecanti Creek Near Sapinero, Co.	71	0.85	1946-72 1911-12, 31-34, 42-99
36	9127500	Crystal Creek Near Maher, Co.	71	0.83	1946-54, 61-69 1911-12, 31-34, 42-99
37	9128500	Smith Fork Near Crawford, Co.	71	0.84	1942-94 1911-12, 31-34, 36-99
38	9129800	Clear Fork Near Ragged Mountain, Co.	71	0.75	1966-73 1911-12, 31-34, 42-99
39	9130500	East Muddy Creek Near Bardine, Co.	71	0.82	1942-53 1911-12, 31-99
40	9130600	West Muddy Creek Near Ragged Mountain, Co.	71	0.82	1956-65 1911-12, 31-34, 42-99

41	9130800	<i>West Muddy Creek Near Bowie, Co.</i>	71	0.69	1969-74 1911-12, 31-34, 42-99
42	9131200	<i>West Muddy Creek Near Somerset, Co.</i>	71	0.72	1962-73 1911-12, 31-34, 42-99
43	9132000	Ruby Anthracite Creek Near Floresta, Co.	71	0.94	1942-43, 55-58 1911-12, 31-34, 39-99
44	9132050	Anthracite Creek Near Somerset, Co.	71	0.92	1977-81 1911-12, 31-34, 42-99
45	9132700	Main Hubbard Creek Near Paonia, Co.	34	0.89	1961-68 1938-99
46	9132800	Middle Hubbard Creek Near Paonia, Co.	57	0.89	1961-68 1959-99
47	9132900	West Hubbard Creek Near Paonia, Co.	34	0.92	1961-73 1938-99
48	9132920	<i>Hubbard Creek Near Bowie, Co.</i>	71	0.71	1969-74 1911-12, 31-34, 42-99
49	9134000	Minnesota Creek Near Paonia, Co.	71	0.9	1942-47, 86-99 1911-12, 31-34, 37-99
50	9134500	Leroux Creek Near Cedaredge, Co.	52	0.91	1939-56, 61-69 1937-99
51	9134700	<i>Cow Creek Near Cedaredge, Co.</i>	52	0.84	1961-69 1939-99
52	9143000	Surface Creek Near Cedared	NA	NA	1940-99
53	9144500	Red Mountain Creek Near Ironton, Co.	34	0.96	1948-56 1938-99
54	9145000	<i>Uncompahgre River At Ouray</i>	NA	NA	
55	9145500	<i>Canyon Creek At Ouray, Co.</i>	71	0.96	1911-12 1911-16, 31- 34, 42-99
56	9146000	<i>Uncompahgre River Below Ou</i>	NA	NA	

57	9146200	Uncompahgre River Near Ri	NA	NA	1959-99
58	9146400	West Fork Dallas Creek Near Ridgway, Co.	34	0.94	1956-70 1938-99
59	9146500	East Fork Dallas Creek Near Ridgway, Co.	57	0.88	1948-53 1959-99
60	9146550	Beaver Creek Near Ridgway, Co.	57	0.69	1961-68 1959-99
61	9146600	Pleasant Valley Creek Near Noel, Co.	71	0.45	1956-67 1911-12, 31-34, 42-99
62	9147000	Dallas Creek Near Ridgway, Co.	71	0.45	1956-71, 80-99 1911-12, 22-27, 31-34, 42-99
63	9147100	Cow Creek Near Ridgway, Co.	34	0.97	1956-73 1938-99
64	9149900	Potter Creek Near Columbine Pass, Co.	52	0.66	1980-81 1939-99
65	9151500	Escalante Creek Near Delta, Co.	71	0.37	1976-89 1911-12, 31-34, 42-99
66	9152000	Kannah Creek Near Whitewater, Co.	52	0.86	1939-82 1918-99
67	9163570	Hay Press C Ab Fruita Res #3, Nr Glade Park, Co.	71	0.71	1984-88 1911-12, 31-34, 42-99
68	9171200	San Miguel River Near Telluride, Co.	34	0.98	1960-65 1938-99
69	9172000	Fall Creek Near Fall Creek, Co.	71	0.97	1942-59 1911-12, 31-34, 41-99
70	9172100	Leopard Creek At Noel, Co.	71	0.71	1956-63 1911-12, 31-34, 42-99
71	9172500	San Miguel River Near Placerville	NA	NA	1911-12, 31-34 1943-99
72	9173500	Horsefly Creek Near Sams, Co.	52	0.53	1943-51 1939-99

73	9174500	<i>Cottonwood Creek Near Nucla, Co.</i>	71	0.64	1942-51 1911-12, 31-34, 42-99
74	9174700	<i>West Naturita C At Upper Sta, Near Norwood, Co.</i>	52	0.79	1975-80 1939-99
75	9176500	<i>Tabeguache Creek Near Nucla, Co.</i>	71	0.58	1946-53 1911-12, 31-34, 42-99
76	9357500	Animas River At Howardsville, Co.	34	0.97	1938-83 1936-99
77	9359000	Mineral Creek Near Silverton, Co.	34	0.97	1938-49 1937-99

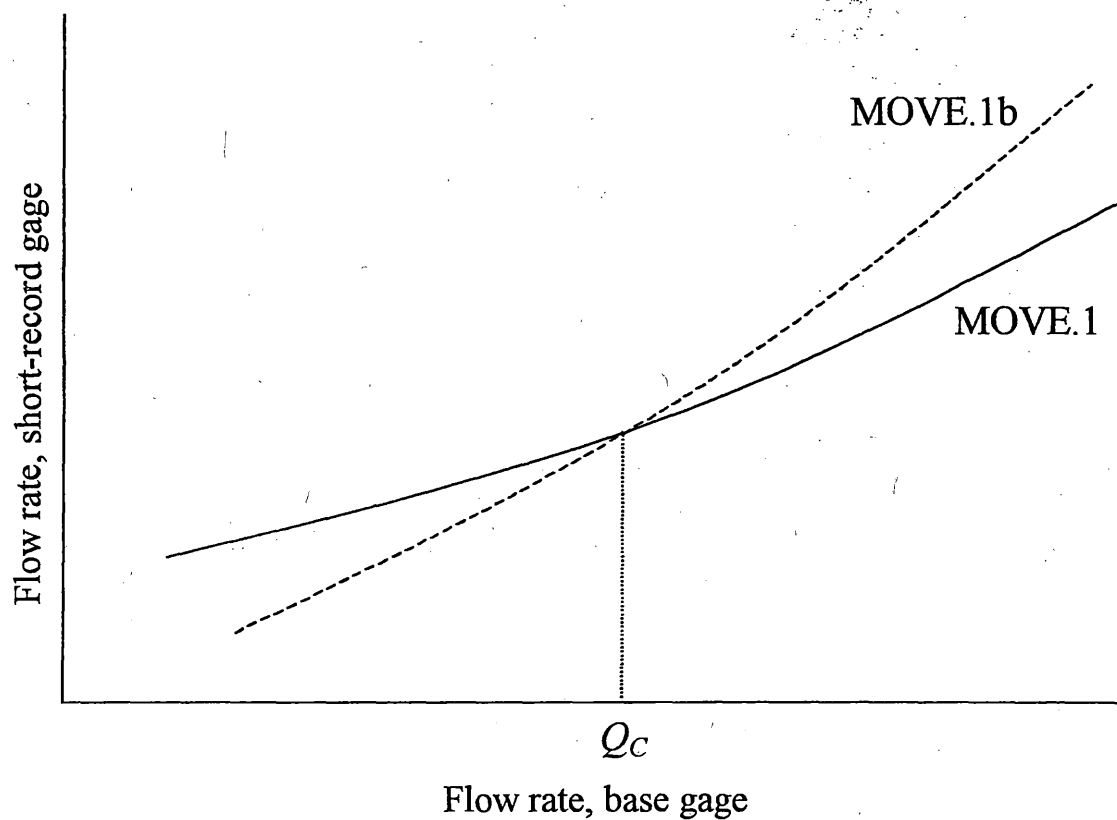


Figure 1 Schematic diagram comparing curves of MOVE.1 and MOVE.1b (mean and variance matched in untransformed space). Their intersection occurs at a base gage flow of Q_C . The relative slopes shown are arbitrary, and the actual difference is generally smaller.

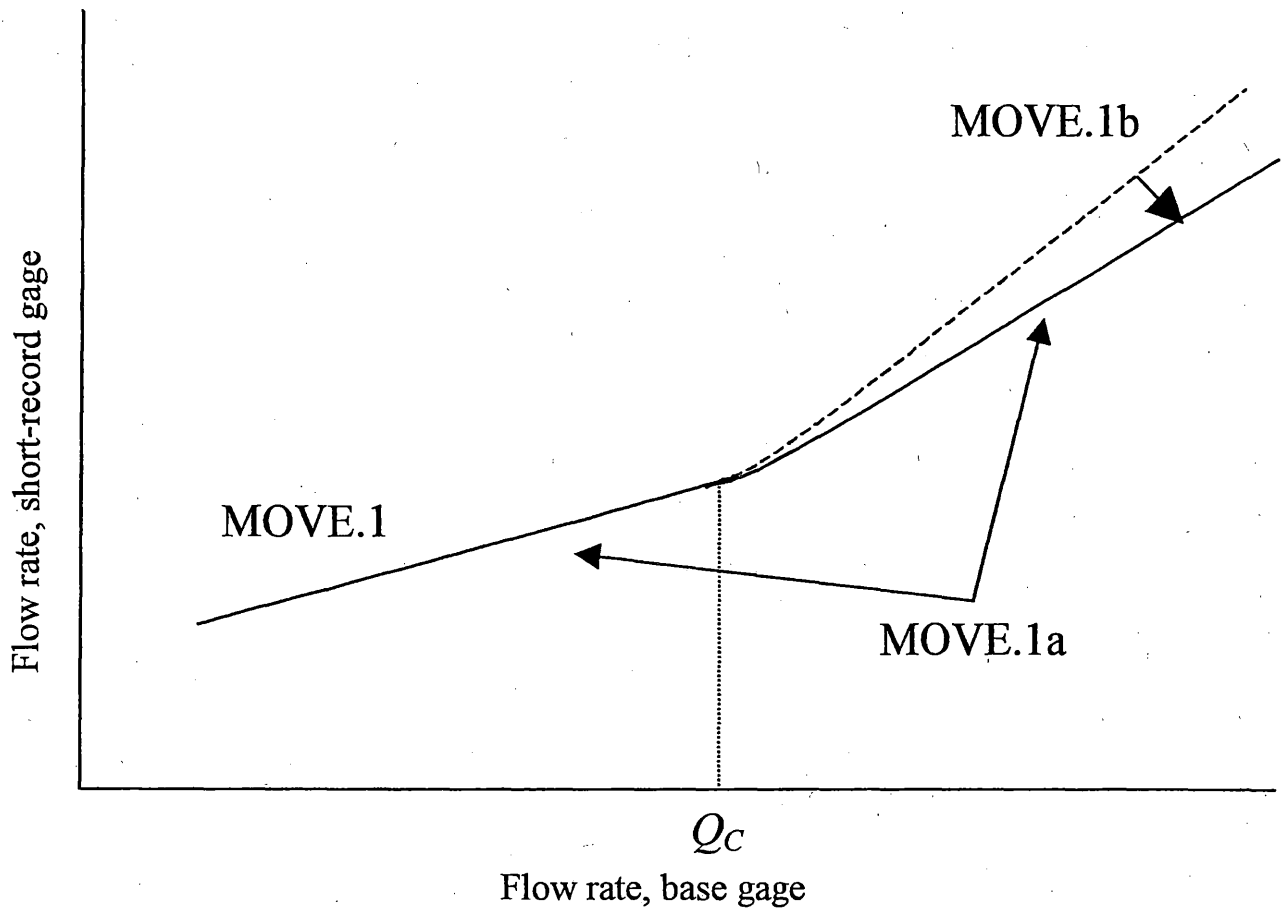


Figure 2 Schematic diagram showing derivation of modified MOVE method, MOVE.1a. The lower branch of MOVE.1 and the upper branch of MOVE.1b form the MOVE.1a curve, after tweaking the upper branch slope (shown here to an exaggerated degree). The branches are divided at the crossover point, Q_C .

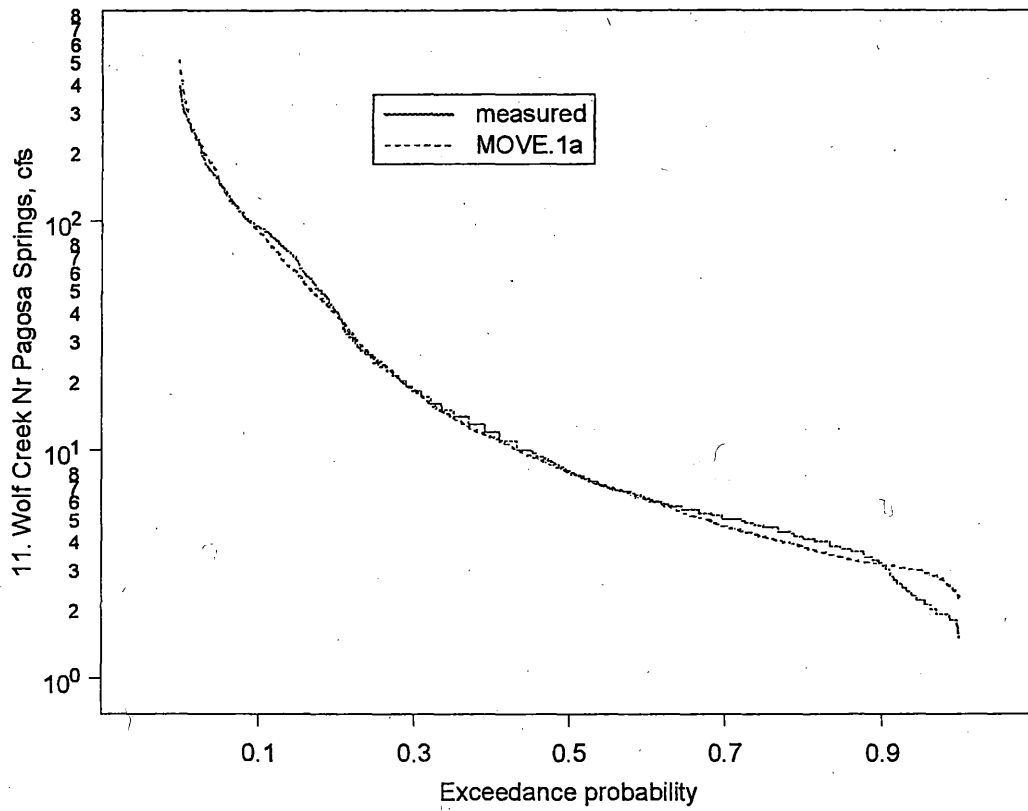


Figure 3 Flow duration curves for the concurrent period, using measured flows and those estimated via modified MOVE.1 (MOVE.1a), for a short record having a correlation 0.95 to its base record, typical for the GMUG National Forest.

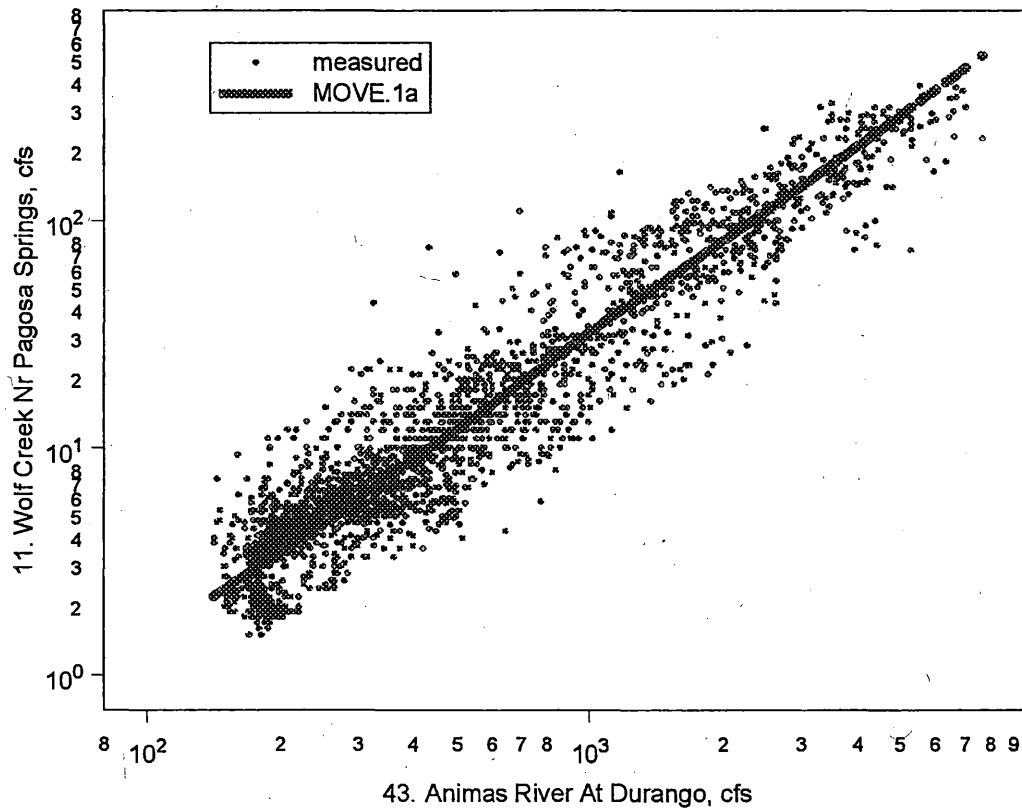


Figure 4 Measured flow rates vs. flow on the same date at the base gage, with the modified MOVE.1 line (MOVE.1a), indicating the estimated flow corresponding to each base station flow. Same gages as Figure 3.

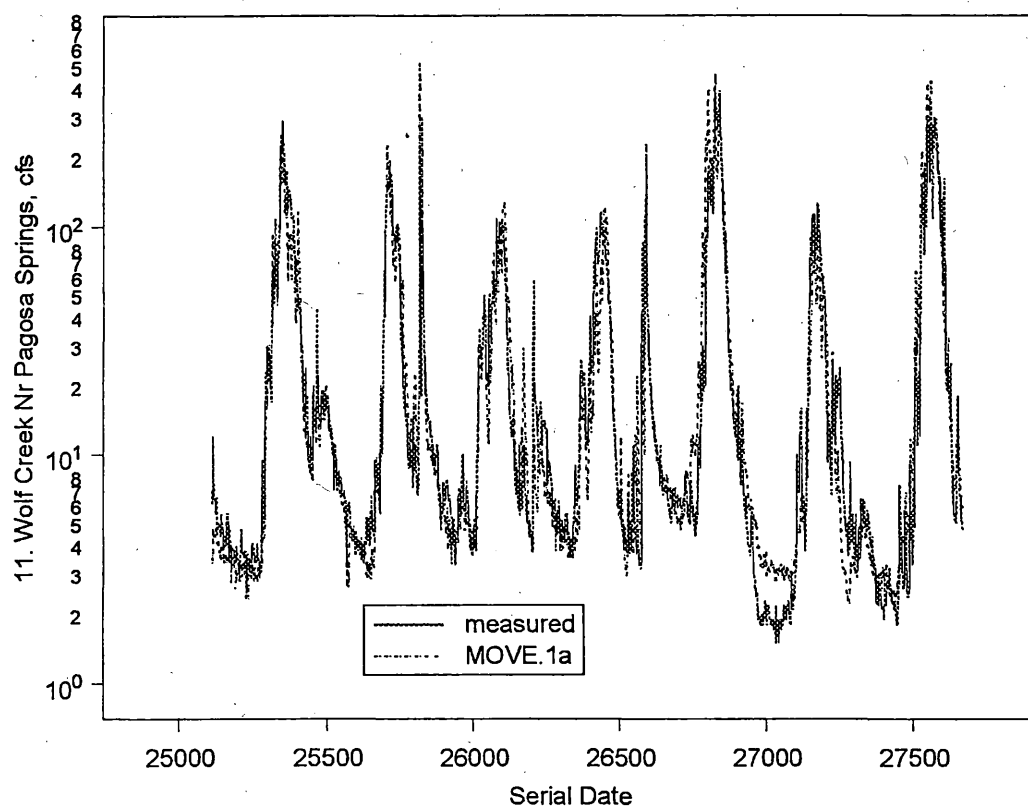


Figure 5 Measured and estimated hydrograph, water years 1969-1975. MOVE.1a is designed to produce a similar mean and variance to the measured record, but not to match specific dates. Same gages as Figure 3.

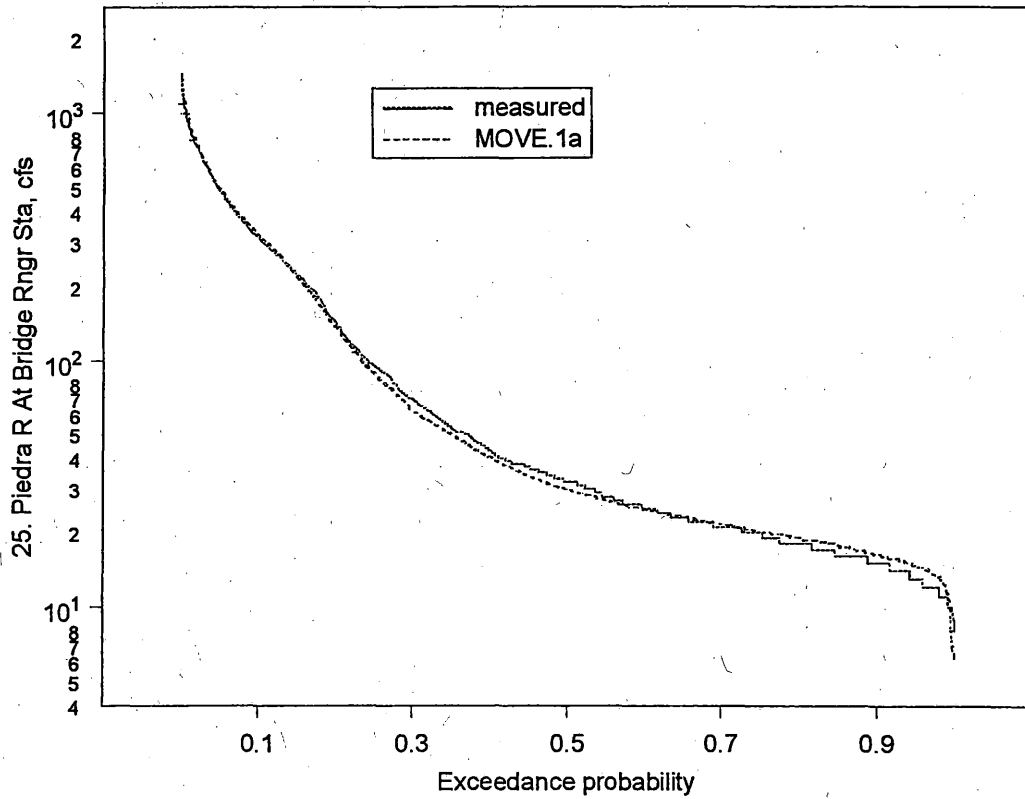


Figure 6 Concurrent-period measured and estimated flow duration curves for a short record which has a high correlation to its base gage (0.97).

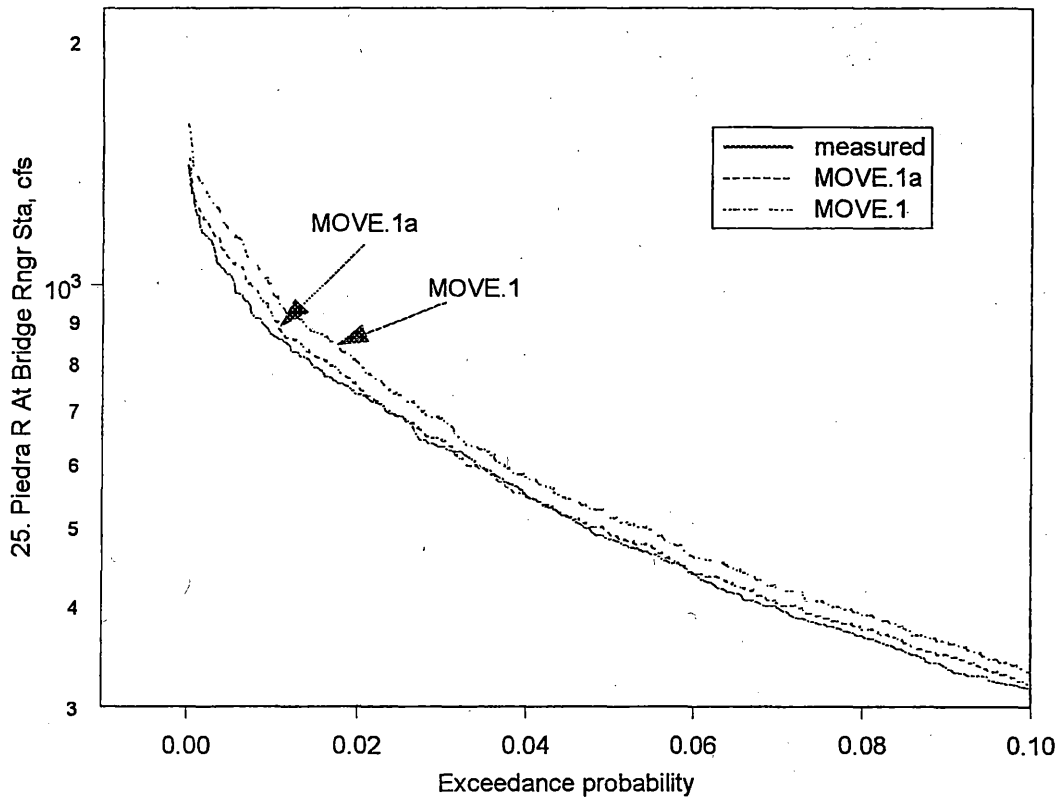


Figure 7 Flow duration curves for the highest 10% of flows in Figure 6, showing the high-flow error associated with MOVE.1 even in a well-correlated pair, and its improvement using the modified method, MOVE.1a.